Introduction to Geomagnetism

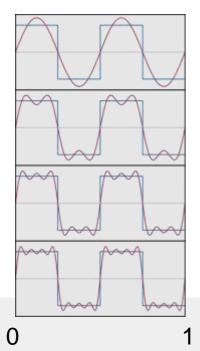
- The magnetic record of Earth's history that is frozen into crustal rocks has provided, perhaps, the most important geophysical evidence of past and on-going plate-tectonic processes.
- The nearly static magnetic field that is our strong dipolar geomagnetic field has been most interesting to geologists and geodynamicists.
- The generation of the field is now the most interesting scientific question for many mathematically oriented geophysicists.
- Why and how we have our field is informative about the processes at work deep within our Earth.



Harmonic functions -- I

We can resolve any function over the line interval [0, 1] in terms of coefficients of a *Fourier Series:*

$$\frac{a_0}{2} + \sum_{n=1}^{\infty} \left[a_n \cos(nx) + b_n \sin(nx) \right]$$



All
$$a_n = 0$$
; $b_2 = 1$

Add
$$b_4 = 1/4$$

Add
$$b_8 = 1/8$$

Add
$$b_{16} = 1/16$$

We can build up our squarewave function coefficient by coefficient. We can construct any function on the interval by adding up the Fourier components scaled by the appropriate coefficient.

Harmonic functions -- II

We resolve the magnetostatic potential at any place over the nearly spherical Earth in terms of coefficients g_i^m and h_i^m of an expansion in **Spherical Harmonics:**

$$P_0^0(\cos\theta) = 1$$

$$P_1^0(\cos\theta) = \cos\theta$$

$$V(r,\phi,\theta) = a \sum_{\ell=1}^{L} \sum_{m=0}^{\ell} \left(\frac{a}{r}\right)^{\ell+1} \left(g_{\ell}^{m} \cos m \phi P_{1}^{0}(\cos \theta) = -\sin \theta\right)$$

$$P_2^0(\cos\theta) = \frac{1}{2}(3\cos^2\theta - 1)$$

Example of the forms of the harmonics:
$$\frac{P_2^1(\cos\theta) = -3\cos\theta\sin\theta}{P_2^1(\cos\theta)} = -3\cos\theta\sin\theta$$

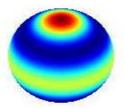
$$g_l^m = 1$$
, $h_l^m = 0$, for $l = 4$, $m = 0$, 1,...4 at $radiv P_2^2(\cos\theta) = 3\sin^2\theta$

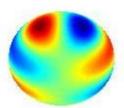
$$P_3^0(\cos\theta) = \frac{1}{2}(5\cos^3\theta - 3\cos\theta)$$

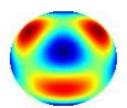
$$m=0$$

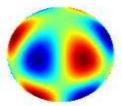
$$m=1$$

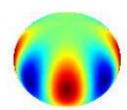
$$m=4$$



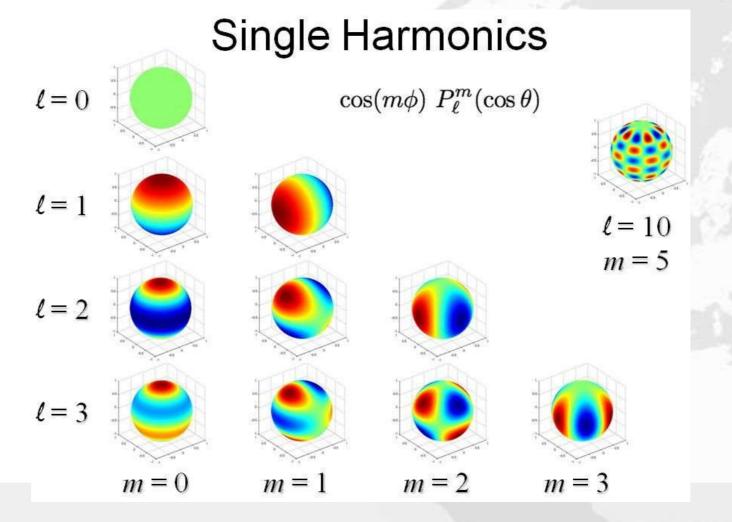








Harmonic functions -- III



Mapping the IGRF

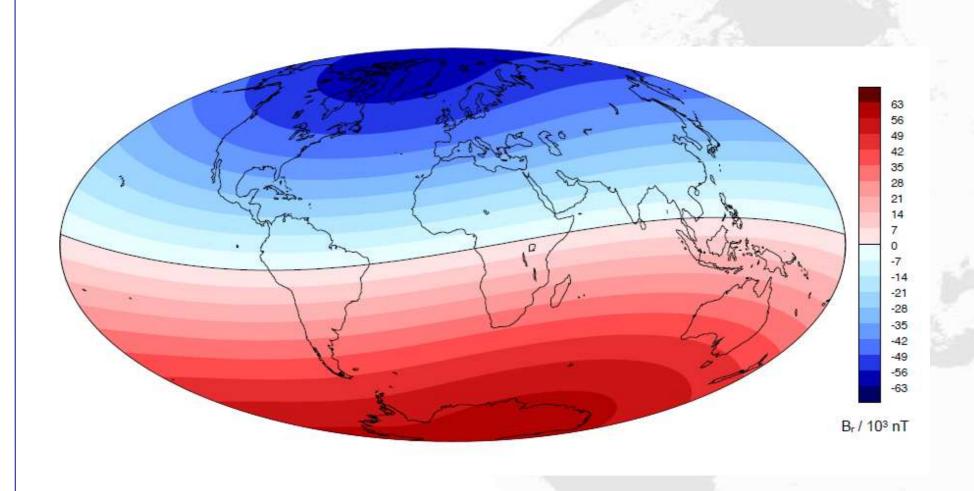
Several satellite missions, most recently, the **persted** and **champ** missions, have mapped the Earth's geomagnetic field from altitude and then downward continued the measurements to a sphere of average Earth radius. The spherical harmonic coefficients obtained from these missions construct the IGRF (International Geomagnetic Reference Field) – that field that is taken as the datum field against which temporal and spatial anomalies are measured.







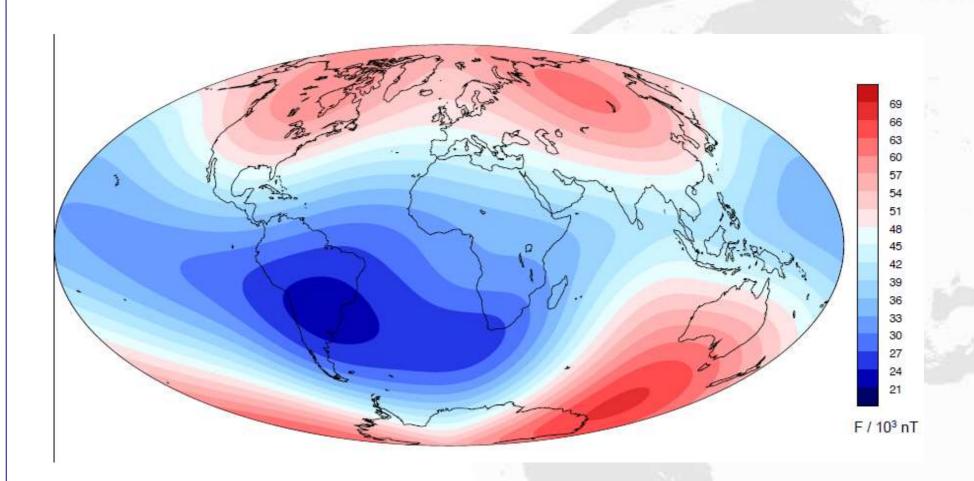
The IGRF - I



The radial (local vertical) component B_z – dipole components only



The IGRF - II



The total field |B| – degree 13 (i.e. all I = 0 to 13 coefficients)

IGRF calculator for any place on the Earth's surface:



Global Magnetic Anomaly – B_z

Residual to the IGRF

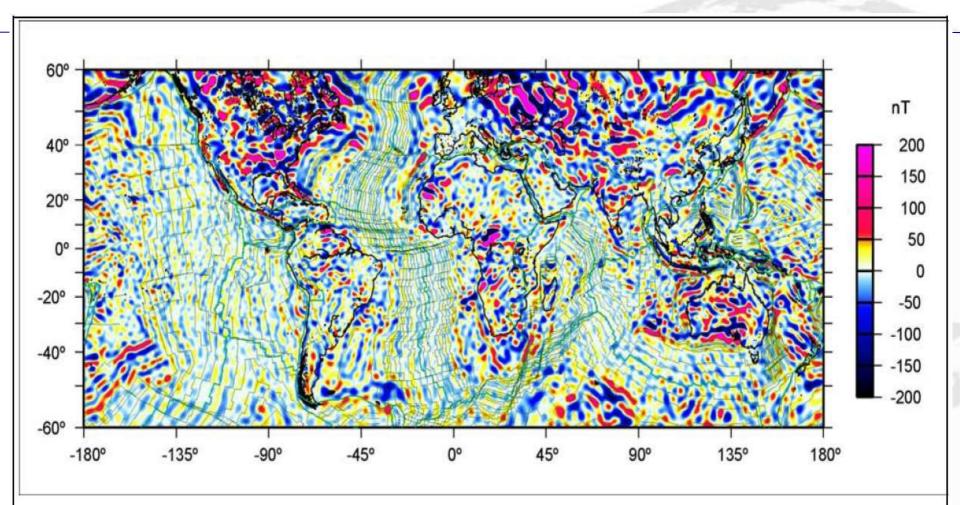


Figure 1: Vertical component of the CHAMP-satellite-derived MF6 crustal magnetic field at the Earth surface, overlain with the isochrons of an ocean-age model inferred from independent marine and aeromagnetic data by Muller et al. (2008).



Elsewhere in the Solar System

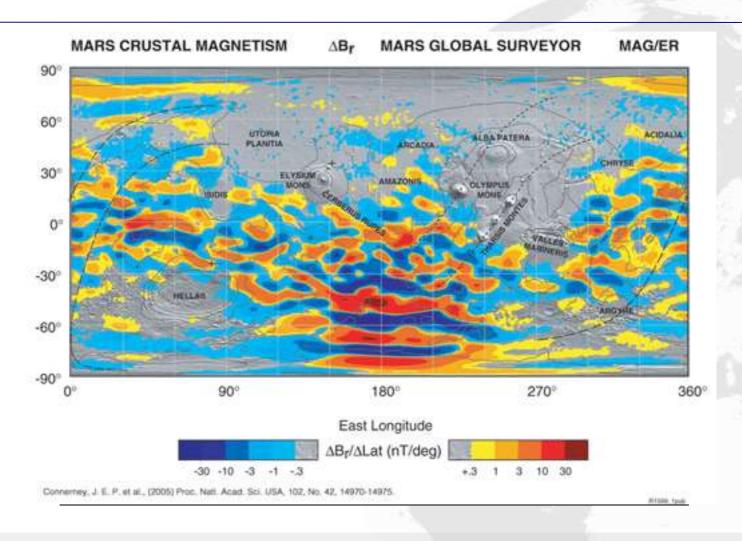
Can we detect a tectonic history written into the crust?

- Most space missions are launched with probes carrying magnetometers.
- The first interest is in the character of the global magnetic fields surrounding planets and moons that might suggest an internal magnetodynamic dynamo.
- Where among the rocky bodies: Mercury, Earth, lo, Ganymede have dynamos.
- Perhaps Mars and Earth's Moon once had dynamos that have imprinted a paleomagnetic record into crustal rocks.



Mars' Magnetic Anomaly

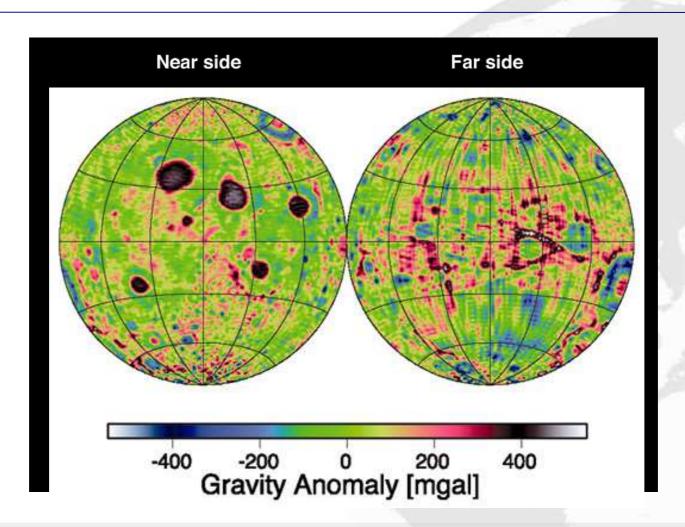
Reveals tectonic history?



Mars Global Surveyor Mission



Lunar Magnetic Anomaly -- |B|



The Apollo astronauts returned rocks from the Moon which showed very high remnant magnetism – <u>Lunar Prospect</u> mapped the field.



Venus' Magnetic Anomaly?

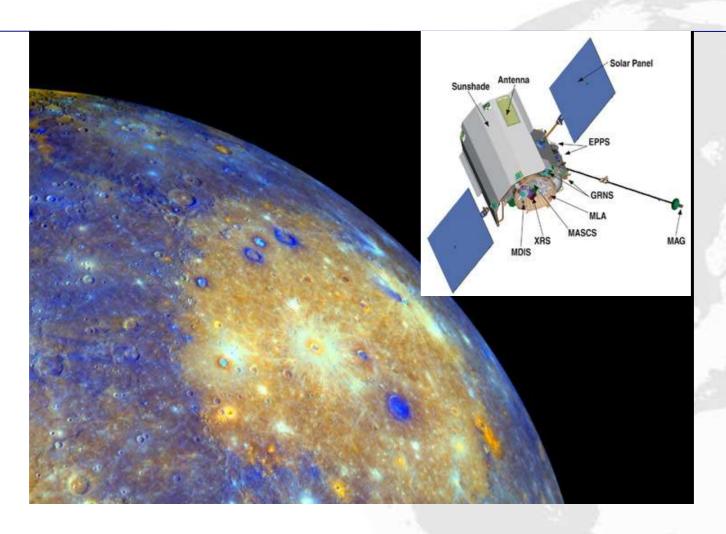
A tectonic history written into the crust?

- No geomagnetic field has been detected on Venus
 probably no convecting core.
- The surface temperature is 740K; most minerals are well above their Curie temperature at 740K – no field imprinted in crustal rocks.
- What we can say about the surface is that it is very "young" -- completely resurfaced within the past 400-700 million years.
- We must infer tectonic history based on surface topography and gravity.



Mercury's Magnetic Anomaly?

The <u>Messenger</u> mission should reveal tectonic history



The story is just now being told!





Plate tectonics

- Morley and Larochelle (1964) And, then, following, Vine and Matthews (1963) recognize paleomagnetic banding across ocean ridges was due to sea-floor spreading.
- During the late 1950s, <u>Edward (Ted) Irving</u> had been mapping paleomagnetic pole paths to show the history of continental drift.
- The age of the sea-floor basins has been established by magnetic surveying since the original recognition of spreading.

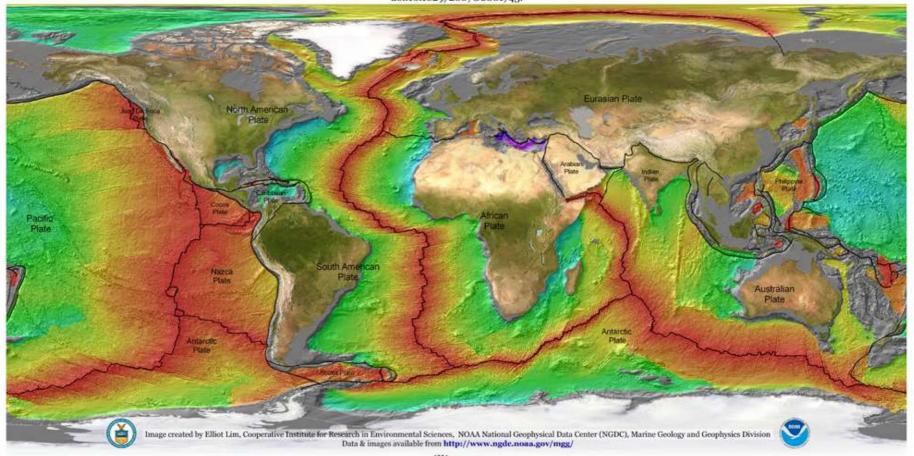


Sea-floor spreading revealed by paleomagnetic anomalies

Age of Oceanic Lithosphere (m.y.)

Data source:

Muller, R.D., M. Sdrolias, C. Gaina, and W.R. Roest 2008. Age, spreading rates and spreading symmetry of the world's ocean crust, Geochem. Geophys. Geosyst., 9, Q04006, doi:10.1029/2007GC001743.



million years
40 60 80 100 120 140 160 180 200 220 240 260 280

Plate tectonics – The theory is demonstrated by paleomagnetic evidence for sea-floor spreading

